# Supplementary information

# Ocean surface energy balance allows a constraint on the sensitivity of precipitation to global warming

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**Supplementary Figure 1. Ocean surface albedo as a function of global mean 2-m temperature.** The annual ocean surface albedo is the ratio of the area-weighted mean outgoing shortwave radiation to the incoming shortwave radiation observed by the Clouds and the Earth's Radiation Energy System (CERES Edition 4.1, <a href="https://ceres.larc.nasa.gov/data/">https://ceres.larc.nasa.gov/data/</a>). Global mean temperature anomalies are from GISS

Surface Temperature Analysis (GISTEMP v4,

<u>https://data.giss.nasa.gov/gistemp/</u>). The solid line represents linear regression with the regression statistics noted (*N*, number of years; *R*, linear correlation coefficient).



Supplementary Figure 2. Relationship between the land modifier ( $\varphi$ ) and tree fraction change. The solid line represents linear regression with the regression statistics noted (N, number of experiments; R, linear correlation coefficient). Error bars are  $\pm$  one standard deviation. Description of model scenarios is given in Figure 1 caption.



## Supplementary Figure 3. Components of the global ocean surface energy

**balance.** The flux is positive in the direction indicated and negative if it goes against the direction shown. a – albedo;  $\beta$  – Bowen ratio; net radiation  $R_n = (1 - a)K_{\downarrow} + L_{\downarrow} - L_{\uparrow}$ .



### Supplementary Figure 4. Interdependence between temperature sensitivities of

### incoming shortwave $(\Delta K_{\perp}/\Delta T)$ and longwave radiation $(\Delta L_{\perp}/\Delta T)$ at the ocean

**surface.** The solid line represents linear regression with the regression statistics noted (N, number of model experiments; R, linear correlation). One outlier, marked by light blue filled triangle with the model name noted, is excluded from the statistical calculation. The white circle with error bars ( $\pm 1$  standard deviation) denotes the observational constraint. The white square denotes sensitivities due to atmospheric moistening under clear skies<sup>1</sup>. Description of model scenarios is given in Figure 1 caption.



Supplementary Figure 5. Comparison between CMIP5 RCP8.5 and CMIP6 ssp585 scenarios. a, Component contributions to global precipitation temperature sensitivity ( $\Delta P/\Delta T$ ). Error bars are  $\pm$  one standard deviation. b, Relationship between changes in global precipitation  $\Delta P$  and ocean evaporation  $\Delta E_0$ , with regression statistics indicated (*N*, number of model experiments; *R*, linear correlation). c, Emergent constraint on global precipitation temperature sensitivity, where the *x*-axis is the same as in Figure 4a. Solid line represents the regression equation in Figure 4a. Description of model scenarios is given in Figure 1 caption.



Supplementary Figure 6. Dependence of precipitation temperature sensitivity  $(\Delta P/\Delta T)$  on feedback strength. Here, the feedback strength  $\alpha$  is approximated by the negative value of temperature sensitivity of ocean heat storage  $(-\Delta G/\Delta T)$  from a  $4 \times CO_2$  simulation using the same models. The solid lines represent linear regression with the regression statistics noted (*R*, linear correlation), Description of model scenarios is given in Figure 1 caption.



# Supplementary Figure 7. Relationship between ocean albedo and ocean incoming shortwave radiation at high latitudes (north of 60° N and south of 60° S). a, annual mean ocean albedo *a* versus annual mean incoming shortwave radiation $K_{\downarrow}$ according to the CERES observation (<u>https://ceres.larc.nasa.gov/data/</u>). b, inter-model spread in the *a* and $K_{\downarrow}$ temperature sensitivities for CMIP5 historical simulations. The solid lines represent linear regression with the regression statistics noted (*N*, number of year in panel a and model experiments in panel b; *R*, linear correlation).



Supplementary Figure 8. Comparison of different definitions of the hydrological climate sensitivity. Data points are annual mean values from the IPSL-CM5A-LR model simulations for three climate scenarios. Solid red circles denote the first and last 10 years of the  $4 \times CO_2$  simulation. Red pluses denote the 10-year mean values. In the present study, the slope of *P* versus *T* is approximated by precipitation temperature sensitivity  $\Delta P/\Delta T$ .



Supplementary Fig. 9. Comparison of regional and global analysis using CMIP5 historical simulations. a, Component contributions to global precipitation temperature sensitivity  $\Delta P/\Delta T$  calculated with Equation (1) using global mean values as inputs. b, Component contributions from a regional diagnostic analysis, where Equation (1) was applied separately to polar (north 60° N and south of 60° S) and nonpolar grids (between 60° N and of 60° S), and the result was weighted by the area fraction of each group to give the global mean value. Red: sum of the five component contributions; yellow: contribution by Bowen ratio change; blue: contribution by surface albedo change; light blue: contribution by change in surface net longwave radiation; grey: contribution by change in ocean heat storage. Error bars are  $\pm$  one standard deviation.



Supplementary Table 1. Regression of changes in global precipitation and in ocean evaporation. For each CMIP scenario, *s* is the slope of linear regression between changes in global precipitation ( $\Delta P$ ) and ocean evaporation ( $\Delta E_0$ ) across models (with intercept forced through zero), where  $\Delta P$  and  $\Delta E_0$  are differences in global precipitation and ocean evaporation, respectively, between the last and the first 10-years of each model simulation. For MERRA-2, *s* is the slope of linear regression between annual global *P* and global  $E_0$  (with intercept forced through zero). Uncertainty range is  $\pm$  one standard deviation, estimated as half of the 95% confidence bound on the regression slope. *N* – number of models (climate scenarios) or number of years (reanalysis);  $\varphi$  – land modifier, the ratio of land evaporation change to ocean evaporation change; *R* – linear regression coefficient. All correlations are significant at *p* < 0.001.

	N	S	φ	R					
Climate model scenario									
CMIP5 Historical	36	$0.754 \pm 0.054$	$0.152{\pm}0.185$	0.987					
CMIP5 RCP2.6	25	$0.822 \pm 0.027$	$0.385{\pm}0.094$	0.986					
CMIP5 RCP4.5	34	$0.821 \pm 0.027$	$0.381{\pm}0.093$	0.966					
CMIP5 RCP6.0	19	$0.825 \pm 0.047$	0.395±0.163	0.946					
CMIP5 RCP8.5	37	$0.783 \pm 0.025$	$0.253{\pm}0.085$	0.952					
CMIP5 4×CO2	25	$0.762 \pm 0.013$	$0.180{\pm}0.046$	0.984					
CMIP6 ssp585	18	$0.842 \pm 0.039$	0.455±0.136	0.962					
Reanalysis									
MERRA-2	38	0.695±0.013	-0.051±0.045	0.992					

Variable	Mean	S.D.	Reference or data source			
Temperature ser	sitivity					
$\frac{\Delta\beta}{\Delta T}$ (K <sup>-1</sup> )	-0.00834	0.000188	Yang & Roderick (ref. 2)			
$\frac{\Delta a}{\Delta T}$ (K <sup>-1</sup> )	-0.00653	0.00147	CERES; Kato et al (ref. 3)			
$\frac{\Delta K_{\downarrow}}{\Delta T} \text{ (W m}^{-2} \text{ K}^{-1}\text{)}$	-2.93	0.276	Reanalysis products			
$\frac{\Delta L_{\downarrow}}{\Delta T} \ (W \text{ m}^{-2} \text{ K}^{-1})$	7.51	0.672	Reanalysis products			
$\frac{\Delta L_{\uparrow}}{\Delta T}$ (W m <sup>-2</sup> K <sup>-1</sup> )	5.24	0.0191	Stefan-Boltzmann Law			
$\frac{\Delta G}{\Delta T}$ (W m <sup>-2</sup> K <sup>-1</sup> )	0.625	0.0257	Cheng et al. (ref. 4)			
Ocean energy balance components						
$R_{\rm n}$ -G (W m <sup>-2</sup> )	116					
$K_{\downarrow}$ (W m <sup>-2</sup> )	185					
a	0.0811		Wild et al. (ref. 5)			
β	0.160					

**Supplementary Table 2. Empirical constraints on the energy balance components at the ocean surface.** Refer to Figure S3 for symbol definitions.

### Supplementary Table 3. Sensitivity of incoming surface shortwave $K_{\perp}$ and

longwave radiation  $L_1$  to global temperature T. The sensitivity value is calculated

as the regression slope of the annual mean  $K_{\downarrow}$  or  $L_{\downarrow}$  over ocean grids against the

global mean temperature and adjusted slightly to remove the bias in  $K_{\perp}$  or  $L_{\perp}$  in reference to the CERES value (<u>https://ceres.larc.nasa.gov/data/</u>). Also shown is the coefficient of determination R<sup>2</sup>. All regressions are significant at p < 0.0001. The

MERRA-2  $\Delta K_{\downarrow}/\Delta T$  (value in parentheses) is excluded from the mean value given in Supplementary Table 2.

	$\Delta K_{\downarrow}/\Delta$	Т	$\Delta L_{\downarrow}/\Delta T$		
Data source	W m <sup>-2</sup> K <sup>-1</sup>	<b>R</b> <sup>2</sup>	W m <sup>-2</sup> K <sup>-1</sup>	R <sup>2</sup>	Period
NOAA-CIRES					1851-
https://psl.noaa.gov/	-2.36	0.27	6.28	0.92	2014
NCEP-NCAR					1948-
https://psl.noaa.gov/	-3.43	0.63	8.53	0.96	2019
JRA-55					1958-
https://rda.ucar.edu/	-3.55	0.63	6.17	0.95	2013
ERA-5					1980-
https://cds.climate.copernicus.eu/	-2.41	0.52	5.54	0.95	2019
MERRA-2					1980-
https://esgf-node.llnl.gov/search/create-ip/	(-9.16)	0.52	8.90	0.89	2019

**Supplementary Table 4. List of CMIP5 and CMIP6 climate model simulations used in this study.** The CMIP5 simulation periods for historical, future (RCP2.6, RCP 4.5, RCP6.0 and RCP8.5) and 4×CO2 scenarios are 1850 - 2005, 2006 - 2100 and 1850 - 1999, respectively. For CMIP6 ssp5-8.5, the simulation period is 2015-2100. Symbol T denotes availability of tree fraction data. Model CMCC-CMS does not output evaporation data and is not used in Fig. 1. BNU-ESM historical experiment is not used for tree fraction analysis presented in Supplementary Fig. 2 because it shows an unrealistically large tree fraction increase in the historical period. Model GISS-E2-R RCP2.6 has an unusually high global precipitation temperature sensitivity. Unless stated otherwise, GISS-E2-R RCP2.6 is excluded from the analysis.

Model name	Historical	RCP2.6	RCP4.5	RCP6.0	<b>RCP8.5</b>	4×CO <sub>2</sub>	ssp585
CMIP5							
ACCESS1.0	Y	Ν	Y	Ν	Y	Y	
ACCESS1.3	Y	Ν	Y	Ν	Y	Y	
BCC-CSM1.1	Y	Y	Y	Y	Y	Y	
BCC-CSM1.1m	Y	Y	Y	Y	Y	Y	
BNU-ESM	Υ, Τ	Υ, Τ	Υ, Τ	Ν	Υ, Τ	Υ, Τ	
CCSM4	Y	Y	Y	Y	Y	Y	
CESM1(CAM5)	Y	Y	Y	Y	Y	Ν	
CMCC-CESM	Υ, Τ	Ν	Ν	Ν	Υ, Τ	Ν	
CMCC-CM	Y	Ν	Y	Ν	Y	Ν	
CMCC-CMS	Y	Ν	Y	Ν	Y	Ν	
CNRM-CM5	Y	Y	Y	Ν	Y	Y	
CSIRO-Mk-3-6-0	Y	Y	Y	Y	Y	Y	
CanESM2	Y	Y	Y	Ν	Y	Y	
FGOALS-g2	Y	Y	Y	Ν	Y	Y	
GFDL-CM3	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	
GFDL-ESM2G	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	
GFDL-ESM2M	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	
GISS-E2-H	Y	Y	Y	Y	Y	Y	
GISS-E2-H-CC	Y	Ν	Y	Ν	Y	Ν	
GISS-E2-R	Y	Y	Y	Y	Y	Y	
GISS-E2-R-CC	Y	Ν	Y	Ν	Y	Ν	
HadGEM2-CC	Υ, Τ	Ν	Υ, Τ	Ν	Υ, Τ	Ν	
HadGEM2-ES	Y	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	
INM-CM4	Υ, Τ	Ν	Υ, Τ	Ν	Υ, Τ	Ν	
IPSL-CM5A-LR	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	
IPSL-CM5A-MR	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	
IPSL-CM5B-LR	Υ, Τ	NT	Υ, Τ	Ν	Υ, Τ	Υ, Τ	
MIROC-ESM	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	
MIROC-ESM- CHEM	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	NT	
MIROC5	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Υ, Τ	Ν	

MPI-ESM-LR	Υ, Τ	Υ, Τ	Υ, Τ	Ν	Υ, Τ	Υ, Τ		
MPI-ESM-MR	Υ, Τ	Υ, Τ	Υ, Τ	Ν	Υ, Τ	Υ, Τ		
MPI-ESM-P	Υ, Τ	Ν	Ν	Ν	Ν	Υ, Τ		
MRI-CGCM3	Y	Y	Y	Y	Y	Y		
MRI-ESM1	Y	Ν	Ν	Ν	Y	Ν		
NorESM1-M	Y	Y	Y	Y	Y	Y		
NorESM1-ME	Y	Y	Y	Y	Y	Ν		
CMIP6								
AWI-CM-1-1-MR							Y	
BCC-CSM2-MR							Y	
CAMS-CSM1-0							Y	
CanESM5							Y	
CESM2							Y	
CESM2-WACCM							Y	
EC-Earth3							Y	
EC-Earth3-Veg							Y	
FGOALS-f3-L							Y	
FGOALS-g3							Y	
GFDL-ESM4							Y	
INM-CM4-8							Y	
INM-CM5-0							Y	
IPSL-CM6A-LR							Y	
MIROC6							Y	
MPI-ESM1-2-HR							Y	
MRI-ESM2-0							Y	
NESM3							Y	

### **Supplementary References**

- 1. Pendergrass, A. G. & Hartmann, D. L. The atmospheric energy constraint on global-mean precipitation change. *J. Clim.* **27**, 757–768 (2014).
- 2. Yang, Y. & Roderick, M. L. Radiation, surface temperature and evaporation over wet surfaces. *Q. J. R. Meteorol. Soc.* **145**, 1118–1129 (2019).
- Kato, S. et al. Surface irradiances of edition 4.0 clouds and the Earth's radiant energy system (CERES) energy balanced and filled (EBAF) data product. *J. Clim.* 31, 4501–4527 (2018).
- 4. Cheng, L. et al. How fast are the oceans warming? *Science* **363**, 128–129 (2019).
- 5. Wild, M. et al. The energy balance over land and oceans: an assessment based on direct observations and CMIP5 climate models. *Clim. Dynam* **44**, 3393–3429 (2015).